Project Final Report

for

A Grant Awarded Through the Section 319(h) Nonpoint Source Implementation Program Cooperative Agreement #C9994861-02

under the

Section 319(h) Kentucky Nonpoint Source Implementation Grant Workplan "Stream Geomorphic Bankfull Regional Curves 2002"

> Kentucky Division of Water NPS 02-05 MOA PO2 0600000744

> > July 1, 2004 to June 30, 2009

UNIVERSITY of LOUISVILLE STREAM INSTITUTE



Benjamin D. Mater Arthur C. Parola, Jr. Chandra Hansen

of

The Stream Institute Department of Civil and Environmental Engineering University of Louisville Louisville, Kentucky

and



Margaret Swisher Jones Kentucky Division of Water Kentucky Energy and Environment Cabinet Frankfort, Kentucky The Energy and Environment Cabinet (EEC) and the University of Louisville Research Foundation, Inc., (a KRS164A.610 corporation) do not discriminate on the basis of race, color, national origin, sex, age, religion, or disability. The EEC and the University of Louisville Research Foundation, Inc., will provide, on request, reasonable accommodations including auxiliary aids and services necessary to afford an individual with a disability an equal opportunity to participate in all services, programs and activities. To request materials in an alternative format, contact the Kentucky Division of Water, 200 Fair Oaks Ln, Frankfort, KY 40601 or call (502) 564-3410 or contact the University of Louisville Research Foundation, Inc.

Funding for this project was provided in part by a grant from the US Environmental Protection Agency (USEPA) through the Kentucky Division of Water, Nonpoint Source Section, to the University of Louisville Research Foundation, Inc., as authorized by the Clean Water Act Amendments of 1987, §319(h) Nonpoint Source Implementation Grant #C9994861-02. Mention of trade names or commercial products, if any, does not constitute endorsement. This document was printed on recycled paper.

Acknowledgements

- Arthur C. Parola, Jr., Ph.D., Professor of Civil and Environmental Engineering (CEE) at the University of Louisville and Director of the Stream Institute (ULSI), was the principal investigator for this project.
- Margi Swisher Jones, Technical Advisor for the Kentucky Nonpoint Source Pollution Control Program, Kentucky Division of Water, provided project oversight.
- Benjamin D. Mater, M.Eng., ULSI Research Project Engineer, directed the collection and analysis of stream gauging and geomorphic data and contributed to the writing of the report.
- William S. Vesely, ULSI Research Project Engineer, assisted with geomorphic and gauge data collection and analysis and developed the bed material-flow velocity relations used in this report.
- Michael A. Croasdaile, Ph.D., Assistant Professor of CEE, assisted with geomorphic data collection and analysis.
- Chandra Hansen, ULSI Research Technical Writer, contributed to the writing of the report and edited the final report.
- D. Joseph Hagerty, Ph.D., Professor of CEE, provided valuable advice regarding geologic descriptions.
- Mark N. French, Ph.D., Ph.D., Professor of CEE, provided valuable advice regarding flood-frequency analysis of gauging stations.
- Nageshwar R. Bhaskar, Ph.D., Professor of CEE, provided valuable advice regarding regression analysis.

David K. Thaemert, M.S., ULSI Research Technologist Senior, assisted with data collection.

- Kelly D. Nickelson, Research Grants Coordinator of the JB Speed School of Engineering, prepared Appendix A.
- Several CEE students contributed to data collection, research, and analysis. C. Davis Murphy, J. Brandon Kolze, Dempsey L. Ballou, and Clayton C. Mastin assisted with geomorphic and sediment data collection and analysis. Hannah R. Gill and J. Duncan Gatenbee assisted with sediment analysis
- The Kentucky Water Science Center of the US Geological Survey (USGS) made possible the use of gauge data in this project.

Contents

Lis	t of Fi	gures and Tables	vi				
Exe	Executive Summary ix						
1	Intro	duction	1				
	1.1	Background	2				
	1.2	Project Purpose and Scope	2				
2	The V	Western Kentucky Coal Field Physiographic Region	3				
	2.1	Structural Geology	3				
	2.2	Physiographic Setting	5				
3	Meas	surement and Analysis Methods	9				
	3.1	Site Selection	9				
	3.2	Data Collection	14				
	3.3	Data Analysis	16				
4	Bank	full Characteristics of Western Kentucky Coal Field Channels	18				
	4.1	Bankfull Regional Curves	18				
	4.2	Bankfull Discharge Recurrence Interval	23				
5	Appl	ication of Bankfull Regional Relations	25				
Ref	ference	es	27				
Ap	pendic	ves	29				
	А	Financial and Administrative Closeout	31				
	В	Quality Assurance Project Plan	33				
	С	Western Kentucky Coal Field Data	43				

Figures and Tables

FIGURES

Figure 2.1	(a) Generalized geologic map of Kentucky(b) Physiographic map of Kentucky	4
Figure 2.2	Geology of the WKCF region	6
Figure 2.3	Gravel-bed stream of southern Hancock County typical of headwater streams of the region's southern, eastern, and western margins (unnamed tributary to West Fork Adams Fork).	8
Figure 2.4	Disturbed silt-bed stream of southern Henderson County typical of lowland streams of the region's interior and northern margin (Beaverdam Creek)	8
Figure 2.5	Silt-bed stream of Hopkins County typical of undisturbed lowland streams of the region's interior and northern margin in which large woody debris and beaver dams influence morphology (Lick Creek)	9
Figure 3.1	Whitelick Creek in southern Henderson County. The bankfull level is represented by a narrow, horizontal depositional bench below and distinct from the higher valley flat.	. 12
Figure 3.2	Well-developed, flat, vegetated bench within a larger disturbed channel along West Fork Adams Fork, Hancock County.	. 13
Figure 3.3	Lick Creek in western Hopkins County. The bankfull level coincides with the valley flat.	. 13
Figure 3.4	Locations of assessment sites	. 15
Figure 4.1	Bankfull cross-sectional characteristics as a function of drainage area for silt-bed streams in the WKCF region	. 20
Figure 4.2	Bankfull cross-sectional characteristics as a function of drainage area for gravel-bed streams in the WKCF region	. 21
Figure 4.3	Bankfull discharge as a function of drainage area for silt- and gravel-bed streams in the WKCF region. Silt-bed streams are shown as solid points; gravel-bed streams are shown as shaded grey points	. 22
Figure 4.4	Regression curves for Q1.01 and Q1.5 derived from WKCF gauges. Bankfull flows estimated for study sites are plotted without regression lines	. 24
Figure 4.5	Bankfull discharge as a proportion of the flow associated with the 1.5-year return interval for the same site. (Drainage area increases from left to right.)	. 25
TABLES		

Table 2.1	WKCF Generalized Stratigraphy	5
Table 3.1	Assessment Site Location Summary	14

Table 3.2	Estimated Average Channel Velocity for Bed Materials in the Southeastern US Coastal Plain	17
Table 3.3	Gauged Sites Used for Flood Frequency Analysis	18
Table 4.1	Bankfull Geometry, Classification, and Discharge Data for Silt- and Gravel-Bed Streams of the WKCF Region	19
Table 4.2	Bankfull Regression Equations for Silt- and Gravel-Bed Streams of the WKCF Physiographic Region	19
Table 4.3	Regression Equations for 1.01- and 1.5-Year Events Based on Gauge Data	23

Executive Summary

Regional geomorphic relations, which describe average bankfull channel geometry and flow as a function of upstream drainage area for streams in a given region, provide a reliable point of reference for assessing stream conditions and evaluating channel dimensions and flow. Data collected from 20 un-gauged sites were used to develop regional curves for bankfull channel depth, width, area, and discharge for rural, unregulated Western Kentucky Coal Field (WKCF) streams draining fewer than 117 mi². The bankfull discharge was also compared to the 1.01- and 1.5-year flows estimated for gauged locations in the WKCF with at least 10 years of record of peak annual flow.

An extensive examination and collection of stream geomorphological characteristics in the WKCF was conducted. Cross sections and longitudinal profiles were surveyed, and flow and bed sediment data were collected to compute bankfull parameters and to identify the channel type according to the Rosgen classification system. The effects of geology, historical land-use, and current land use on sediment loads and channel evolution were also considered in stream assessment and in development of the curves.

Bankfull regional curves were derived from collected data by using ordinary least-squares regression to relate bankfull channel dimensions and estimates of bankfull discharge to drainage area. The relationship between bankfull parameters and drainage area in the WKCF region is described by curves that explain between 54% and 92% of the variation within the datasets for bankfull area, width, depth, and discharge. Standard errors are less than 36%. At each of the sites for which bankfull discharge was estimated, the bankfull discharge was found to be no more than 15% of the estimated regional (WKCF) 1.5-year discharge and is thus likely associated with a smaller return interval. Therefore, use of the 1.5-year event for bankfull flow would represent a gross overestimate in most of the reaches examined in this study; estimates based on morphological features and/or the regional curves would be more accurate.

Geomorphic Characteristics of Streams in the Western Kentucky Coal Field Physiographic Region of Kentucky

By Benjamin D. Mater, Arthur C. Parola, Jr., Chandra Hansen, and Margaret Swisher Jones

1. Introduction

The physical characteristics of stream channels strongly influence aquatic and riparian habitat, bank erosion, and sediment loads. Siltation, habitat modification, and flow alteration are the cause of nearly half of the identified stream impairments in the Commonwealth (KDOW 2008), with siltation cited most frequently. These primarily physical causes of stream impairment are all dependent on the presence of riparian vegetation; the entrainment, transport, and storage of sediment; and other geomorphic characteristics of stream channel networks. Changes in these characteristics are a product of complex watershed processes and human modification of the watershed and the stream channel network. Disturbance of streams due to land-use practices such as development, livestock grazing, land clearing, road construction, and channel modification or relocation tend to increase stream peak flow rates, disturb riparian vegetation, and alter stream channel characteristics. The response of many streams to disturbance can be excessive production of sediments through channel incision and subsequent severe bank erosion. The disturbances of streams and the associated erosion that continues for long periods can severely degrade stream and riparian habitat not only at the disturbed section of the stream but upstream and downstream as well.

To determine the implications of various physical impacts, specific geomorphic data are needed to evaluate flow stresses, sedimentation, and other physical habitat factors affecting biological communities. In assessing channel stability and habitat, estimates of bankfull flow conditions are particularly useful for

- Classification of the stream reach using the Rosgen (1996) method
- Determination of the degree to which the stream is incised
- Indication of relative bank stability
- Indication of some characteristics of channel pattern
- Indication of the capacity of the channel to transport its supplied load

Evaluation of channel stability is essential for the assessment of sediment loads, which may be needed for development of the sediment total daily maximum loads (TMDLs) required by recent US Environmental Protection Agency guidelines (USEPA 1999). Moreover, geomorphic data from a watershed's streams, including estimates of bankfull parameters, can be used as a basis for the design of stream restorations which physically alter disturbed stream channels in order to

improve stream habitat, reduce bank erosion, and reduce sediment loads. In designing stream restorations, estimates of bankfull characteristics are useful for

- Initial estimation of channel geometry for planning of a restoration project prior to detailed morphological assessments required for final design
- Estimation of channel design parameters for sites where morphological characteristics are inconsistent or have not been developed
- Comparison of restoration designs developed using other methods
- Evaluation of restoration designs by permit agencies

Estimates of bankfull parameters may be obtained through direct measurement of similar channels in a watershed or region, or they may be obtained through analytical procedures such as the development of an effective discharge. They may also be obtained through the use of regional curves, which describe average bankfull width, depth, cross-sectional area, and discharge as a function of upstream drainage area for streams in a given region. Given the strong influence of local climate and geology on stream channel form, regional curves are typically developed with respect to physiographic region (e.g., Brush 1961; Harman et al. 1999; Kilpatrick and Barnes 1964; Leopold et al. 1964; McCandless and Everett 2002; Smith and Turrini-Smith 1999; Wolman 1955). While regional curves do not account for all sources of variability in channel characteristics, their formulation does include consideration of geologic conditions, land use, and valley use, and they provide a reliable point of reference for assessing stream conditions and evaluating channel dimensions and flow.

1.1 BACKGROUND

Quantitative geomorphology has been used for over half a century to support the assessment of channels and floodprone areas. Hydraulic geometry relations developed by Leopold and Maddock (1953) described the relationship between channel dimensions and mean annual discharge within specific drainage basins. In the next decade, hydraulic geometry relations, or at-a-station curves, were developed for several geographic regions in the eastern US (Brush 1961; Kilpatrick and Barnes 1964; Leopold et al. 1964; Wolman 1955). After the introduction and deliberation of the concept of a bankfull discharge, whose stage is just contained within the stream banks (Wolman and Leopold 1957; Wolman and Miller 1960), bankfull channel geometry and discharge data were also collected in the early 1970s (Emmett 2004). In the late 1970s, Dunne and Leopold (1978:614) noted the correlation between bankfull channel parameters and drainage area, and they introduced curves describing average bankfull channel dimensions and bankfull discharge as a function of drainage area in "hydrologically homogenous" regions. In the last decade, regional curves have been developed for physiographic regions across much of the US (e.g., NRCS 2007).

1.2 PROJECT PURPOSE AND SCOPE

Stream restoration efforts intended to improve stream habitat have been conducted in the Western Kentucky Coal Field (WKCF) region without general information on the geomorphic characteristics of streams in various regions of the state. The main purpose of this project was to provide quantitative descriptions (regional curves) that would represent expected values and variation of bankfull flow and channel cross-sectional area, width, and depth as a function of upstream drainage area in the WKCF physiographic region of Kentucky.

Data used to develop the regional curves were collected from 20 sites on 17 un-gauged WKCF streams in March and April of 2007; bankfull discharge was estimated for 11 of these sites. While the inclusion of channels with active stream-flow gauges was a priority in order to be able to estimate the bankfull flow return period, none were suitable for bankfull geomorphic assessment; therefore, data were collected only on un-gauged streams. Criteria used to identify suitable stable stream channels for regional curve data collection included a wide range of drainage basin areas within the physiographic region and as many streams as possible having a channel environment that was alluvial, relatively stable, and showed no signs of ongoing rapid morphological change (cf. McCandless and Everett 2002; Smith and Turrini-Smith 1999). Bankfull regional curves for silt-bed streams draining from 1.56 to 117 mi² and gravel-bed streams draining from 0.25 to 2.88 mi² were derived from the data by using ordinary least-squares regression to relate bankfull discharge and bankfull channel dimensions to drainage area.

2. The Western Kentucky Coal Field Physiographic Region

The WKCF region encompasses all or part of 20 counties and covers roughly 4,800 mi². Those counties that lie almost entirely within the region are Daviess, Hancock, Henderson, Hopkins, McLean, Muhlenberg, Ohio, Union, and Webster. Portions of Butler, Edmonson, and Grayson Counties comprise its eastern extension. The major stream systems of the region include the Ohio, the Green, the Tradewater, the Pond, the Rough, and the Nolin rivers.

The physiographic regions of Kentucky correspond to geologic regions of the state (Figures 2.1a and b), as the effects of surface weathering and erosion of different geologies produce landscapes and streams of dissimilar characteristics. In the Western Kentucky Coal Field, gravelbed streams dissect the Pennsylvanian rocks of the uplands, while silt-bed streams rework Quaternary sediments of the lowlands.

2.1 STRUCTURAL GEOLOGY

The Western Kentucky Coal Field physiographic region is part of a larger physiographic region known as the Illinois Basin, which extends throughout southeastern Illinois, southwestern Indiana, and into western Kentucky. The WKCF is bounded to the north and northwest by the Ohio River and to the west, south, and east by the Dripping Springs Escarpment, formed from sandstones and conglomerates in the lower part of the Pennsylvanian strata. WKCF Pennsylvanian rocks (Table 2.1 and Figure 2.2) are roughly 4000 ft thick and occupy the broad, subtle Moorman Syncline. The east-west trending axis of this syncline parallels and is 7-10 mi south of the region's most prominent band of normal and reverse faults: the Rough Creek Fault System (Rice 2001). Along the region's southern and eastern margin, basal Pennsylvanian strata meet underlying strata of Mississippian limestone to form the Pottsville Escarpment. The resistant Pennsylvanian rock forms a dissected plateau within the WKCF that resembles the physiography of the Eastern Kentucky Coal Field (EKCF) region. The similarity of portions of the two regions is due in part to their locations on the flanks of the Cincinnati Arch (Burroughs 1924), where formations have eroded less rapidly than those nearer its axis, which is oriented in an approximately north-south direction between Cincinnati, Ohio, and Lexington, Kentucky. The Axis formed in a series of folding and warping episodes that lifted Paleozoic strata far above the elevations at which they had been deposited (McFarlan 1943) and now separates the Pennsylvanian strata that were once continuous between the WKCF and EKCF regions.



Figure 2.1 (a) Generalized geologic map of Kentucky (after McGrain 1983:12). (b) Physiographic map of Kentucky (KGS 1980).

During the Pennsylvanian period (250-300 mya), sediment eroding from the ancestral Appalachian hills was deposited in a large inland sea extending throughout the Illinois Basin. Fluctuations in the level of this ancient sea, along with basin subsidence and changes in depositional environment, resulted in a cyclical layering of the region's coal-bearing lithology, comprised predominantly of interbedded sandstone, shale, coal, and to a lesser extent, limestone. These largely alluvial or deltaic strata are of similar origin and characteristics to those of the Eastern Kentucky Coal Field. The massive, quartzose sandstone of the WKCF Caseyville and Tradewater Formations, for example, resembles that of the Lee Formation in the EKCF. In both regions, the erosion-resistant quality of this rock type is responsible for its presence in prominent cliff outcroppings. Overlying this resistant sandstone unit in the WKCF are the interbedded sandstones, siltstones, coal, and limestones of the Tradewater, Carbondale, and Sturgis formations. Clastic rock types dominate these lithologically similar formations, with marine carbonates forming less prominent but extensive layers formed by intermittent sea transgression. Economically valuable coal seams are most prominent within the middle of the Pennsylvanian stratigraphy, particularly within the Carbondale Formation (Rice 2001).

System	Series	Formation	Description
	Upper Pennsylvani- an	Sturgis	Sandstone, siltstone, shale, limestone, and coal: Sandstone, light-gray and light- to yellowish-brown, fine- to coarse-grained. Siltstone, light- to dark-gray, locally shaly and sandy. Shale, medium- to dark-grey, carbonaceous, generally silty and sandy. Limestone, gray, thin- to thick-bedded, fossiliferous. Coal is generally a minor constituent but locally very thick.
/ANIAN	msylvanian	Carbondale	Sandstone, siltstone, shale, coal, and limestone: Sandstone, light- to medium- gray or brown, very fine- to medium-grained, medium- to thick-bedded; Silt- stone, medium-gray, sandy; Shale, light-brown to dark-gray, sandy, silty, partly limonitic and carbonaceous. Many thicker and more economic coal beds of the WKCF.
DENNSYLV	Middle Pen	Tradewater	Sandstone, siltstone, shale, coal, and limestone: Sandstone, light- to medium- gray, fine- to medium-grained; micaceous, friable. Siltstone, light- to medium- gray, locally interbedded with thin-bedded sandstone. Shale, light-gray to black, carbonaceous. Several economic coal beds. Limestone, light- to medium-gray, dense, thin- to thick-bedded, fossiliferous.
	Lower Pennsylvanian	Caseyville	Sandstone, siltstone, shale, and coal: Sandstone, white to yellowish-brown, fine- to coarse-grained; locally conglomeratic with scattered quartz pebbles as much as 2 inch in diameter; locally forms cliffs as much as 125 feet high. Siltstone, light- to dark-gray; locally micaceous and interbedded with thin-bedded sand- stone. Shale, dark-gray to black, carbonaceous. Coal beds are thin, lenticular, and of very local extent.

Table 2.1 WKCF Generalized Stratigraphy*

* The typical stratigraphy of the region provided by the US Geological Survey (Hansen and Smith 1978; Palmer 1968; Amos 1970; Rice 2001).

2.2 PHYSIOGRAPHIC SETTING

The physiograpy of the WKCF ranges from dissected uplands with V-shaped valleys in the south and east and the western margin to gently rolling hills and wide, flat bottomlands in the north and interior areas.

Dissected Uplands

The region's southern and eastern margin is a dissected plateau with narrow ridges and valleys. Within this portion of the region, the close proximity of Pennsylvanian stratigraphy leads to its direct interaction with surficial geomorphic processes. As in the EKCF, this interaction has created a heavily dissected plateau drained by a dendritic pattern of mostly gravel-bed streams, although regional uplift has not provided the same expanse of rugged terrain as in the EKCF. Stream morphology within the relatively high-relief WKCF areas bears some resemblance to that found in similar areas of the EKCF (Figure 2.3).

The western extent of the WKCF is similar in character, though not as rugged, as the southern and eastern margins. The Tradewater River runs north-west along the boundary between the dissected uplands of the western extent and the rolling hills of the interior and north. Tributaries draining from the west to the Tradewater in Crittenden and Caldwell Counties encounter Mississippian material at headwater extremities and may locally encounter outcroppings of durable conglomerate sandstones of the Caseyville formation. Rock types along these tributaries may transition sharply where faulting has occurred. Headwater tributaries above widespread lacustrine and alluvial deposits of the lowlands often have gravel beds with high sand content. These smaller streams resemble those in the eastern portion of the WKCF and in some areas of the EKCF.



Figure 2.2 Geology of the WKCF region (KGS 2002; KYDGI 2005b; Noger 2002).

In the transition zone between these uplands and the more subdued topography of the north and interior, stream morphology is more variable due to changes in sediment regimes and boundary material. The transition is abrupt where upland tributaries encounter slackwater deposits of larger rivers as with many streams draining to the Tradewater River in Crittenden and Caldwell Counties.

Rolling Hills and Wide Valley Bottomlands

From the rough rim of Christian, southern Muhlenberg, Butler, Edmonson, Grayson, and eastern Ohio counties, the topography transitions to gently rolling hills within central Hopkins, Muhlenberg, and Ohio counties. Fine-grained sediment and alluvium derived from loess deposits are common and play a significant role in stream morphology. Along the Green River, alluvium is up to 100 ft thick, and depths of over 50 ft are common (Palmer 1972). Tributaries draining from the east to the Tradewater River traverse Quaternary lacustrine and alluvial deposits and are thus similar to those feeding the Green River on the east of the Green/Tradewater drainage divide (Franklin 1969). Lowland streams are low–gradient and silt-bedded, and where they have

not been modified, they form multiple channel streams flowing through wetlands in wide valley bottoms filled with Quaternary silt, sand, and gravely alluvium.

Topographic relief becomes even more subdued in the interior region consisting of Webster, northern Hopkins, northern Muhlenberg, McLean, and northwestern Ohio counties. Within this area, loess is more common and mantles all but the steepest slopes to depths up to 25 ft (Hansen and Smith 1978). This loess is generally a soft, clayey silt and is easily eroded, thus constituting a major portion of sediment supplied to the area's streams. Below elevations of approximately 400 ft msl, clay and silt of valley bottoms is often of lacustrine origin, deposited in Quaternary lakes formed as a result of Wisconsin outwash in the Ohio River Valley (Johnson and Smith 1972). Alluvium found in the valley bottoms is chiefly clay and silt with deposits of sandy gravel. Alluvial gravels are mostly subangular to subrounded pebbles of chert with smaller quantities of rounded pebbles of quartz, sandstone, siltstone, and fossiliferous limestone. Gravels are often cemented in limonitic deposits (Hansen and Smith 1978). Anthropogenic disturbance is widespread in these areas and propagates quickly throughout drainage networks, while morphological recovery appears to be slow (Figure 2.4). Although some valley bottoms are swampy and have stream morphology heavily influenced by large woody debris and beaver dams (Figure 2.5), most have been extensively drained for agriculture (Burroughs 1924). Streams within this area are generally bedded in silt with minor coarse bedload and are of low gradient. Headwater tributaries may have gravel beds associated with reworked Pleistocene and, potentially, Pliocene deposits (Hansen and Smith 1978) and eroded bedrock at headward extents.

Gently rolling loess-mantled physiography extends into the WKCF's northern counties of Union, Henderson, and Daviess. At the northern extent of the region, a distinct transition from rolling upland to river bottom parallels the Ohio River, becoming most prominent in the steep bluffs and narrow tributary valleys of Hancock County in the northeast. Loess deposits are up to 55 ft thick, and the larger streams occupy wide valley bottoms filled with deep deposits of Pleistocene and Holocene alluvium (Johnson 1973). Alluvium is sand, silt, gravel, and clay derived from contemporary deposition and older glacial outwash. Gravels are mostly subangular to subrounded quartz and chert pebbles less than 30 mm in length with larger particles at greater depths (Johnson 1973). Like the streams of the region's interior, streams of the northern counties transport mostly clay, silt, and sand-sized sediment and have beds and banks formed of the same material. Gravel bedload seems to be only a minor component of the total sediment load and thus has little observed influence on stream morphology with the exception of small headwater tributaries. Bottomlands that have not been drained are generally swampy, and backwater effects from the Ohio River may play an important role in the morphology of many streams. Morphological impacts of drainage enhancement projects such as channelization and debris removal are widespread, and such perturbations propagate swiftly throughout stream networks. As in the interior region, morphological recovery appears to be slow.

The mantling of fine-grained material derived from Pleistocene loess and thick deposits of alluvial and lacustrine material characteristic of the northern and interior portions of the WKCF are ubiquitous in the nearby Mississippi Embayment region. In both regions, this mostly fine-grained material dominates the sediment load of lowland streams, and through its accretion and erosional characteristics it plays a major role as a control on channel morphology. The similarities in sediment and topographic relief suggest that the low-gradient, silt-bed streams within this portion of the WKCF may resemble those found in the Mississippi Embayment.



Figure 2.3 Gravel-bed stream of southern Hancock County typical of headwater streams of the region's southern, eastern, and western margins (unnamed tributary to West Fork Adams Fork).



Figure 2.4 Disturbed silt-bed stream of southern Henderson County typical of lowland streams of the region's interior and northern margin (Beaverdam Creek).



Figure 2.5 Silt-bed stream of Hopkins County typical of undisturbed lowland streams of the region's interior and northern margin in which large woody debris and beaver dams influence morphology (Lick Creek).

3. Measurement and Analysis Methods

Bankfull channel characteristics were measured at un-gauged sites throughout the WKCF region. Channel cross-section, longitudinal profile, and bed material data were used to calculate channel dimensions and parameters needed for classifying the channel, developing bankfull regional curves, and at some locations, estimating bankfull discharge. The bankfull discharge was also compared to the 1.01- and 1.5-year flows estimated for gauged locations in the WKCF with at least 10 years of record of peak annual flow.

3.1 SITE SELECTION

Initial Screening of USGS Gauging Stations

When bankfull conditions can be identified at gauging stations, discharge can be related to the bankfull stage, and a frequency can be estimated for the bankfull flow event. Therefore, USGS gauging stations with drainage areas less than 200 mi² were initially considered in the selection of a sample to represent the region's population of streams. In order for the sample to be useful in the development of bankfull regional curves, it would ideally consist of sites on rural, unregulated, wadeable streams with active gauges and a wide distribution of drainage areas and geographic locations. Prior data collection in other physiographic regions of Kentucky, however, had shown that the number of gauge sites suitable for assessment is typically limited and unlikely to comprise a sample that meets all of the ideal criteria; channel conditions at stream gauge stations tend to be characterized by reach-scale instability, a lack of consistent and unambiguous

bankfull indicators in incised channels, and recently modified channel geometry (Parola, Skinner, et al. 2005; Parola, Vesely, et al. 2005). Therefore, while geographic locations and drainage areas were identified and recorded, their distributions were not factors in site selection.

Twelve stations were screened, and half of those were eliminated according to three preliminary selection criteria prior to field reconnaissance:

- Recording frequency and duration of available discharge data. Discontinued gauge sites were excluded unless the record of annual maximum series data was suitable for flood frequency analysis. At least 10 years of data had to be available, spanning a period where the only breaks in the record were those unrelated to flood magnitude (USIACWD 1982:15). Active gauge sites with fewer than 10 years of annual maximum series data were excluded unless they had real-time discharge data for estimating bankfull flow.
- 2. Land use. Because streams in watersheds with a significant proportion of densely urbanized land tend to be undergoing rapid morphological change, watersheds that were more than 10% urbanized and those known to be undergoing urbanization were excluded. Extensive drainage projects and channelization have occurred in all of the watersheds in the region; therefore, associated land uses were not used as a basis for exclusion of sites.
- 3. Site characteristics. Sites known to have characteristics that would make them unsuitable for data collection (e.g., those that were known to be regulated, affected by waterway structures, or undergoing rapid morphological change) were excluded.

Contour maps and aerial photographs were then reviewed to identify characteristics that could be relevant to field evaluation of the sites that had not been eliminated from consideration. The following tasks were completed in the review:

- 1. All stations were located on 1950s or later revisions of USGS 7.5-minute topographic quadrangle maps.
 - a. Reaches likely to present consistent and reliable bankfull indicators were identified.
 - b. Stream reaches in the vicinity of the gauges were examined for evidence of channel straightening, realignment, or other modifications.
 - c. Abandoned, remnant channels were identified.
 - d. Any structures spanning or encroaching on the stream channel were identified.
 - e. Valley constrictions or sharp bends that could create backwater during high flows were recorded.
- 2. Aerial photographs were examined to identify land use changes and possible impacts to the stream channel and the floodplain that had occurred since the creation of the topographic and geologic maps.
- 3. The bedrock material underlying each site and its watershed were identified from Kentucky Geological Survey (KGS) 7.5-minute geologic quadrangle maps.
- 4. Surface drainage areas for each station were recorded from the total drainage areas provided with USGS gauge descriptions. Field reconnaissance was limited to streams in watersheds draining fewer than 200 mi².
- 5. The boundaries of the watershed of each station were identified using geospatial datasets (KGS 2002; Noger 2002). None of the streams draining fewer than 200 mi² had significant portions of their watersheds outside the WKCF region.

Initial Screening of Un-Gauged Sites

The limited number of potential gauged sites necessitated that un-gauged sites be included in the sample selected to represent the region's population of streams. Un-gauged candidate sites on active channels and on remnant channels bypassed by channelized reaches were identified on the USGS 7.5-minute quadrangle maps; sinuous reaches that appeared to be un-channelized were screened according to land use and site characteristics (see above). The un-gauged sites' drainage areas and geographic locations were also included as factors in screening and selection in order to produce a distribution broadly representative of the regional stream population. Because streams draining fewer than 10 mi² are the focus of the majority of natural channel design efforts (i.e., those that would make use of regional curves), their representation in the sample was considered a priority. Approximately 35 un-gauged sites were selected for further evaluation.

Field Reconnaissance

An initial reconnaissance visit was made to photograph and evaluate each potential gauged and un-gauged site. Channel conditions throughout the region, as in other physiographic regions of Kentucky (Parola, Skinner, et al. 2005; Parola, Vesely, et al. 2005), were frequently found to be characterized by reach-scale instability, a lack of consistent and unambiguous bankfull indicators in incised channels, and recently modified channel geometry. The field evaluation eliminated each of the remaining six gauged sites and all but 20 un-gauged sites based on four criteria:

- 1. Access. To obtain morphological data, a stable reach had to be accessible. Sites on private land were only selected if landowners granted access.
- 2. Channel pattern. Only single-thread channels were selected.
- 3. Channel morphology. Sites that met the above criteria were given further consideration only if the channel showed no signs of ongoing rapid morphological change and the geomorphic characteristics of the reach were suitable for surveying of bankfull indicators.
- 4. At gauged sites, physical gauge configuration was capable of recording flows within the range of that which may be considered bankfull; gauges that recorded higher flows only were eliminated. A site was also rejected if the estimated bankfull flow was not within the range of flow data used in developing the gauge rating curve.

The suitability of the channel for surveying of bankfull indicators was determined based on evaluation of the floodplain and channel morphology upstream and downstream of the gauge or within the un-gauged reach. At a minimum, the reach had to have (1) cross-sectional geometry with unambiguous indicators of the bankfull level and evidence of at least one bank having been formed by deposition, (2) channel geometry that was not controlled by a structure, and for gauged sites, (3) a drainage area that differed by no more that 10% from the drainage area at the gauge station. The bankfull level was determined according to the definition of bankfull flow proposed by Dunne and Leopold (1978), who described it as the flow that completely fills the channel so that its surface is level with the active floodplain. The active floodplain is the flat depositional surface adjacent to the channel that is constructed by the present river in the present climate and is frequently inundated by the river (Dunne and Leopold 1978). Dunne and Leopold also reported an approximately 1.5-year average return interval for bankfull flow; in the identification of the active floodplain of WKCF assessment reaches, however, no minimum or maximum bankfull return period was assumed.

The primary indicators used to identify the active or actively-forming floodplain were finegrained depositional features (Dunne and Leopold 1978). The characteristics of these features varied depending on channel morphology. Many incised channels had multiple depositional surfaces—flat terraces that had to be distinguished from the active floodplain. In those channels, the primary indicator was a low depositional bench, and the bankfull level was identified as the point at which the slope transitioned between steep and horizontal (Figure 3.1). In cases where smaller, indistinct channels were forming within an incised channel, a primarily flat, vegetated bench was the most consistently observed depositional feature (Figure 3.2). Other incised channels lacked flat terraces; instead, the region between the valley flat and the channel was only a gently sloped incline, often with active accretion of fine sediment, and the bankfull level was not identifiable. In streams that were not incised, the bankfull level coincided with the top of bank and valley flat (Figure 3.3). Some of these non-incised channels were remnants (Figure 3.4) bypassed by channelized reaches. In many cases, these remnants were functioning as sloughs or slackwater areas of intense deposition of fine-grained material; therefore, only the distance between the tops-ofbank was considered to be representative of bankfull channel geometry.

Identification of the bankfull level was refined by comparing elevations of multiple indicators and evaluating secondary, non-morphological indicators. The elevations of bankfull indicators along the channel were compared to confirm that they were consistent relative to the water surface. When consistent indicators suggested a number of possible bankfull levels, the reach was nevertheless considered to be suitable for surveying. Secondary indicators of the bankfull level included the size fraction of the depositional material and changes in vegetation above and below the level identified as bankfull.



Figure 3.1 Whitelick Creek in southern Henderson County. The bankfull level is represented by a narrow, horizontal depositional bench below and distinct from the higher valley flat.



Figure 3.2 Well-developed, flat, vegetated bench within a larger disturbed channel along West Fork Adams Fork, Hancock County.



Figure 3.3 Lick Creek in western Hopkins County. The bankfull level coincides with the valley flat.

Final Site Selection

A total of 20 un-gauged reaches on 17 streams were selected as study sites; 12 reaches were on active channels, and 8 were on remnant channels (Table 3.1). Drainage areas ranged from 0.25 mi^2 to 117 mi². No gauged reaches were selected.

Table 3.1	Assessment Site Location Summary	
-----------	----------------------------------	--

	Stream Site	Drainage Area (mi ²)	County	Latitude	Longitude
1	Caney Creek*	117.00	Ohio	N37.460317	W86.653767
2	Drakes Creek	29.10	Hopkins	N37.174000	W87.444083
3	Drakes Creek at Old Nortonville-Whiteplains Road	31.79	Hopkins	N37.180783	W87.435267
4	Eagle Creek Tributary	0.77	Union	N37.637350	W87.903333
5	East Fork of Flynn Fork	2.88	Caldwell	N37.140817	W87.758800
6	Hazel Creek*	10.97	Muhlenberg	N37.147150	W86.978667
7	Lewis Creek	8.30	Ohio	N37.350500	W86.912117
8	Lick Creek	19.66	Hopkins	N37.268317	W87.714917
9	Lick Creek at Paul Peyton Rd	23.00	Hopkins	N37.284317	W87.723033
10	Muddy Creek*	83.15	Butler	N37.169600	W86.772783
11	No Creek*	9.40	Rough	N37.487933	W86.987983
12	Otter Creek	40.36	Hopkins	N37.489533	W87.384417
13	Pup Creek*	26.13	Daviess	N37.842100	W86.968533
14	Slover Ditch Tributary*	1.56	Webster	N37.447383	W87.729333
15	Welch Creek*	45.47	Butler	N37.261217	W86.595100
16	West Fork Adams	0.87	Hancock	N37.688833	W86.690250
17	West Fork Adams near Newton Springs Church	1.53	Hancock	N37.682867	W86.694717
18	West Fork Adams Tributary	0.25	Hancock	N37.687983	W86.692667
19	West Fork Pond River*	43.00	Christian	N37.108244	W87.363503
20	Whitelick Creek	3.82	Henderson	N37.639967	W87.700800

* Remnant channel.

3.2 DATA COLLECTION

At all sites, sufficient channel and overbank topographic data and bed sediment data were collected to calculate bankfull parameters and to identify the channel type according to the Rosgen (1996) classification system.

Channel Geometry

During field reconnaissance on un-gauged streams, assessment reaches were identified and morphologic data was subsequently collected for the development of regional curves. Extensive photographic documentation was recorded for all sites. The specific geomorphic features that were recorded included bankfull indicators, bed configuration, bed material, bank condition, flow patterns, valley configuration, dominant vegetation, and any structures that might affect flow within the channel or over the valley bottom. Field surveys recorded cross-sectional and, at some sites, longitudinal profile data.



Figure 3.4 Locations of assessment sites. All blue line streams having a drainage area of more than 2 mi² are shown (KGS 2002; KYDGI 2005a, 2005b, 2006).

Survey data were collected according to the procedures described in Harrelson et al. (1994). Survey control points consisting of at least two wooden stakes or steel re-bar were installed at each cross section location. Where practical, cross sections and longitudinal profiles were surveyed using a Nikon DTM-352 total station; measurements were accurate to within 1 cm in both the horizontal and vertical directions. Collected survey data were stored on a hand-held data log-ger during field activities and then transferred to a spreadsheet software program for analysis. Where a total station survey was not practical due to the remoteness of the site location, no longitudinal profile was surveyed, and cross section data were collected with the use of a line level, station tape, and elevation rod according to the following procedures:

- 1. A string line was pulled taut across the channel between the cross-section control points and set to a level horizontal plane using a line level. The taut line established a horizontal datum to which all elevation data were referenced.
- 2. A graduated survey tape was strung parallel to the datum line with zero set at the left bank (facing downstream).
- 3. The cross section was measured by noting the cross-channel station and measuring the vertical distance from the datum using a graduated tape. Distance, elevation, and point description data were recorded by a second survey crew member.

All cross sections were surveyed at locations that both coincided with a clear bankfull indicator and were representative of the reach morphology: at the crest of a riffle whenever possible or, at sites where no well-developed riffle was located in a reach with clear bankfull indicators, at a plane-bed section of the longitudinal profile. In reaches where multiple cross sections were taken, the cross section taken at the most clearly defined riffle crest was used to compute bankfull parameters. Selection of the most appropriate riffle crest for computing bankfull parameters was based on an extensive examination of the reach and its bankfull indicators. Only after the bankfull level was determined was the most appropriate riffle crest selected for surveying. Cross sections were surveyed to the width of the floodprone area or, when the floodprone width was clearly greater than four times the bankfull width, to a point at least one bankfull width from the top of each bank. Longitudinal profiles measured the elevations of the thalweg, water surface, bankfull indicators, and top of bank at several locations along the assessment reach.

The amount and extent of the survey data collection at each site depended largely on site accessibility and whether the study reach was an active or a remnant channel. Where accessibility allowed for transport of conventional survey equipment to the site, multiple cross sections and the longitudinal profile were surveyed. At remote sites where a total station survey was not feasible, only cross sections were surveyed. At one site where the channel had recently incised and at sites in those remnant channels functioning as sloughs or slackwater areas of intense deposition of fine-grained material, only bankfull width was measured.

Bed Sediment Characteristics

The surface particle-size distribution was evaluated at each site on the riffle or plane-bed location surveyed to compute bankfull parameters. The size class corresponding to the median sediment size was visually estimated. For classification purposes, this estimate was considered to be representative of the dominant particle size throughout the reach.

3.3 DATA ANALYSIS

Cross Sections and Profiles

Survey data were reduced using AutoCad. Cross section and longitudinal profile data were then extracted from AutoCad and plotted using Microsoft Excel. Each surveyed cross section at each site was plotted at a 1:1 horizontal-to-vertical scale so that breaks in slope could be clearly identified. Based on each cross section plot, multiple parameters were analyzed as follows:

- Bankfull indicators on both banks were identified and evaluated on each cross section plot to confirm that they corresponded to the active floodplain. Where bankfull indicators suggested a number of possible bankfull levels, the level indicated by the lowest depositional features that were consistent relative to the water surface elevation was selected as bankfull. Where un-incised reaches exhibited natural levee formation, bankfull dimensions were calculated with reference to the consistent valley flat elevation beyond the levees.
- The cross section taken at the most clearly defined riffle crest or plane-bed reach at each site was used to compute bankfull parameters needed for
 - Developing regional curves: bankfull cross-sectional area (A_{BKF}); bankfull width (W_{BKF}); and mean bankfull depth ($D_{BKF} = A_{BKF} / W_{BKF}$).
 - ^a Classifying each assessment reach according to the Rosgen (1996) Level II classification system: maximum bankfull depth; floodprone width (W_{FP}); entrenchment ratio (ER = W_{FP} / W_{BKF}); and width-to-depth ratio (W_{BKF} / D_{BKF}).

 Cross section plots were compared to photographs of the same locations. Banks and depositional features in each cross section plot were examined in the photographs to evaluate their stability. For sites where flow would be estimated, the dominant size fraction of the bed sediment was identified from the photographs for use in assigning average bankfull velocities.

The longitudinal profile of each surveyed channel thalweg, water surface, bankfull indicators, and top-of-bank elevation were plotted with an exaggerated vertical scale so that breaks in slope could be clearly identified. The locations of cross sections were also plotted on each longitudinal profile. Based on each profile plot, bankfull levels at each cross section location were verified. A regression line was plotted through elevations of all bankfull indicators that were consistent relative to the water surface elevation. Where bankfull indicators suggested a number of possible bankfull levels, the level indicated by the lowest depositional features that were consistent relative to the water surface elevation was selected as bankfull. The regression line represented the average bankfull level through the reach. The bankfull level indicated by the regression line was then used to re-evaluate cross section plots: where a residual for a bankfull level point at a cross section was large, or where no bankfull indicator elevation was plotted, the corresponding cross section plot was examined to determine whether a bankfull indicator could be identified close to the level indicated by the regression line.

Bankfull Discharge

Bankfull discharge at 10 active channel sites and 1 remnant channel site was estimated from measured cross-sectional area and assumed average bankfull velocity. Velocity estimates were based on average channel velocities developed for streams of different bed material size classes in the Southeastern US Coastal Plain (Table 3.2). The error associated with the use of these average velocities in discharge calculations is assumed to be no greater than that associated with other less practical methods that would require extensive surveying and numerical modeling.

Bed Material	Average Channel Velocity at Bankfull Conditions (feet per second)
Silt	1.5
Sand	2.0
Bimodal gravel and sand	2.5
Gravel	3.5

Table 3.2	Estimated Average Channel Velocity for Bed Materials in the
Southeastern	ו US Coastal Plain (Parola, Vesely, et al. 2005)

Flood Flow Frequency at Gauging Stations

Annual maximum series data for the eight USGS gauging stations in the WKCF region with more than 10 years of record (Table 3.3) were obtained from the USGS Kentucky Water Science Center or from their online datasets. Using the log-Pearson Type III distribution (McCuen 1998) as described by USIACWD (1982), frequency analysis was conducted for each of the eight stations. From the frequency distribution, flows corresponding to the 1.01- and the 1.5-year events were estimated for each station.

USGS Site Name	Gauge ID	County	No. Yrs. Q _{peak} Data*	Drainage Area (mi ²)	Latitude	Longitude
1 W Fk Adams Fk Tributary near Fordsville	03319520	Hancock	11	0.26	N37.686944	W86.692500
2 Rhodes Creek Tributary near Owensboro	03321465	Daviess	10	0.29	N37.786389	W87.206111
3 Rose Creek at Nebo	03384000	Hopkins	30	2.10	N37.382778	W87.633056
4 Beaverdam Creek near Corydon	03322360	Henderson	19	14.30	N37.703889	W87.697778
5 Bear Branch near Leitchfield	03312000	Grayson	34	30.80	N37.426667	W86.279167
6 South Fork Panther Creek near Whitesville	03321350	Ohio	15	58.20	N37.618889	W86.887500
7 Caney Creek near Horse Branch	03318800	Ohio	36	124.00	N37.463889	W86.655556
8 Pond River near Apex	03320500	Muhlenburg	65	194.00	N37.122222	W87.319444

 Table 3.3
 Gauged Sites Used for Flood Frequency Analysis

* The number of years (through water-year 2007 or, in the case of discontinued gauges, the last year of recorded data) for which (1) peak data was available online from the USGS and (2) the only breaks in the record were those unrelated to flood magnitude.

4. Bankfull Characteristics of Western Kentucky Coal Field Channels

Bankfull channel parameters calculated for each assessment reach (Table 4.1) were used to develop regional curves for WKCF streams and to classify each reach. The curves describe the relationships between drainage area and bankfull channel geometry and bankfull discharge. Estimates of bankfull discharge for 11 sites were also compared to estimates of the regional 1.01-and 1.5-year discharges. Classification of each reach according to Rosgen (1996) Type II classification parameters identified 3 Bc-, 2 C-, and 15 E-type channels; the stream type was consistent for the entire length of each reach.

4.1 BANKFULL REGIONAL CURVES

Bankfull regional curves for silt-bed streams draining from 1.56 to 117 mi² and gravel-bed streams draining from 0.25 to 2.88 mi² were derived using ordinary least-squares regression. Bankfull channel geometry and discharge data were plotted as a function of drainage area on a log-log scale (Figures 4.1-4.3). A best-fit line was regressed for each plot in the form of a simple power function:

$$Y_{bkf} = a DA^b$$
(4.1)

where *a* and *b* are empirically-derived constants, DA is drainage area (mi²), and Y_{bkf} represents a bankfull channel parameter: cross-sectional area, A_{bkf} (ft²); width, W_{bkf} (ft); mean depth, D_{bkf} (ft); or discharge, Q_{bkf} (cfs). The resulting regression equations are provided in Table 4.2 along with calculated coefficients of determination and standard errors.

Coefficient of determination (\mathbb{R}^2) values show that drainage area accounts for over 75% of the variation in the relationships between drainage area and channel bankfull parameters for gravel-bed streams. In silt-bed streams, drainage area accounts for over 54% of the variation. Variation unaccounted for by drainage area may be attributed to other influences such as variability in sediment load caliber and quantity, hydrology, and the effects of local controls (Knighton 1987). For example, the relatively weak relationship between drainage area and the bankfull parameters of silt-bed streams may reflect the influence of beaver dams and large woody debris. Such local controls are common in these low-gradient streams, but their influence

on channel morphology is highly variable, depending on relative debris size, in-channel jam configuration, degree of beaver activity, and degree of channel disturbance.

	Total							Rosgen		
Stream Site	DA (mi ²)	A _{bkf} (ft ²)	W _{bkf} (ft)	D _{bkf} (ft)	ER*	W/D Ratio	Dominant Substrate	Stream Type [†]	Q _{bkf} (cfs)	Q1.5 (cfs) [‡]
1 Caney Creek [§]	117.00	237.3	46.5	5.10	>2.2	9.12	Silt	E6	356	5200
2 Drakes Creek	29.10	216.4	39.0	5.55	>2.2	7.03	Silt	E6	325	2100
3 Drakes Creek at Old Norton- ville-Whiteplains Road	31.79		31.8				Silt	E6		
4 Eagle Creek Tributary	0.77	8.3	12.8	0.65	1.4	19.66	Gravel	C4	25	200
5 East Fork of Flynn Fork	2.88	14.2	13.6	1.04	1.6	13.07	Gravel	B4c	43	500
6 Hazel Creek [§]	10.97		17.7				Silt	E6**		
7 Lewis Creek	8.30	57.5	23.4	2.45	>2.2	9.56	Silt	E6	86	900
8 Lick Creek	19.66	71.0	32.4	2.19	>2.2	14.76	Silt	E6	107	1600
9 Lick Creek at Paul Peyton Rd	23.00	126.4	36.1	3.50	>2.2	10.29	Silt	E6	190	1800
10 Muddy Creek [§]	83.15		35.1				Silt	E6**		
11 No Creek [§]	9.40		36.4				Silt	E6**		
12 Otter Creek ^{††}	40.36	147.4	31.5	4.68	9.1	6.73	Silt	E6	221	2600
13 Pup Creek [§]	26.13		39.3				Silt	E6**		
14 Slover Ditch Tributary [§]	1.56		13.0				Silt	E6**		
15 Welch Creek [§]	45.47		32.0				Silt	E6**		
16 West Fork Adams	0.87	9.1	10.9	0.83	1.6	13.20	Gravel	B4c	27	200
17 West Fork Adams near Newton Springs Church	1.53	14.8	16.2	0.92	1.4	17.66	Gravel	B4c	44	300
18 West Fork Adams Tributary	0.25	3.6	6.0	0.60	2.4	10.00	Gravel	C4	11	100
19 West Fork Pond River [§]	43.00		48.0				Silt	E6**		
20 Whitelick Creek ^{‡‡}	3.82		23.3				Silt	E6**		

Table 4.1 Bankfull Geometry, Classification, and Discharge Data for Silt- and Gravel-Bed Streams of the WKCF Region

* ER is entrenchment ratio (dimensionless).

† Visually estimated slopes were used for classification purposes.

‡ Estimated from WKCF regional regression (see Table 4.3).

§ Remnant channel.

** Probable stream type before abandonment or incision.

†† The study site was approximately 3.4 miles upstream of a confluence with a larger stream (Pond River), but bankfull geometry appeared to be consistent with study sites not influenced by a downstream confluence.

‡‡Recently incised; only bankfull width measured.

Bed Material	Regression Equation	n	Coefficient of Determination, R ²	Standard Error*, Se (%)
	$A_{bkf} = 19.31 DA^{0.56}$	6	0.72	35.1
	$W_{bkf}\ =\ 14.43\ DA^{0.25}$	15	0.66	21.8
Silt bed	$D_{bkf} = 1.23 \ DA^{0.33}$	6	0.54	30.4
	$Q_{bkf} \;\; = \; 28.97 DA^{0.56}$	6	0.72	35.1
	$A_{bkf} = 9.27 DA^{0.60}$	5	0.92	19.2
Crowel had	$W_{bkf}\ =\ 11.56\ DA^{0.36}$	5	0.75	22.0
Gravel bed	$D_{bkf} \ = \ 0.80 \ DA^{0.24}$	5	0.86	10.0
	$Q_{bkf} \;\; = \; 27.80 DA^{0.60}$	5	0.92	19.2

 Table 4.2
 Bankfull Regression Equations for Silt- and Gravel-Bed Streams of the WKCF Physiographic Region

* Transformed from the log10 domain as a percentage of the mean according to Tasker (1978).



Figure 4.1 Bankfull cross-sectional characteristics as a function of drainage area for silt-bed streams in the WKCF region.



Figure 4.2 Bankfull cross-sectional characteristics as a function of drainage area for gravel-bed streams in the WKCF region.



Figure 4.3 Bankfull discharge as a function of drainage area for silt- and gravel-bed streams in the WKCF region. Silt-bed streams are shown as solid points; gravel-bed streams are shown as shaded grey points.

4.2 BANKFULL DISCHARGE RECURRENCE INTERVAL

Because bankfull geometry data could be collected only at un-gauged sites, no annual peak flow data for the study sites were available for use in flood frequency analysis. Likewise, even though several WKCF gauging stations had flow data suitable for flood frequency analysis (Section 3.3), none of the gauged channel reaches were suitable for geomorphic assessment, and flow at the gauging stations could not be related to the bankfull stage. Therefore, bankfull return periods were not calculated for the study sites. Bankfull discharge of the un-gauged study sites were, however, compared to the regional 1.01- and the 1.5-year flows estimated from WKCF gauging station data.

Regional curves for flows corresponding to the 1.01- and the 1.5-year events were derived using ordinary least-squares regression. Estimates of flows corresponding to the 1.01- and the 1.5-year events were plotted as a function of drainage area on a log-log scale for the eight WKCF gauge stations for which frequency analysis was conducted (Figure 4.4). Best-fit lines were regressed in the form of a simple power function:

$$Q_{t-return} = a DA^{b}$$
(4.2)

where *a* and *b* are empirically-derived constants, DA is drainage area (mi²), and $Q_{t-return}$ is the discharge with a return interval of *t* years. The resulting regression equations are provided in Table 4.3 along with calculated coefficients of determination and standard errors.

Table 4.3	Regression Equations for 1.01- ar	10 I.5-Y	ear Events Based on G	auge Data
			Coefficient of	Standard Error*,
Region	Regression Equation	Ν	Determination, R ²	S e (%)
WKCF	$Q_{1.01} = 88.81 \text{ DA}^{0.69}$	8	0.94	53.9
	$Q_{1.5} = 236.85 \ DA^{0.65}$	8	0.95	44.5

Fable 4.3 Regression Equations for 1.01- and 1.5-Year Events Based on Gauge Data

* Transformed from the log10 domain as a percentage of the mean according to Tasker (1978).

The bankfull discharges that were estimated for 11 study sites were plotted to compare them to the regression lines. Bankfull discharge at each of the study sites was less than the estimated regional 1.01-year discharge for the same drainage areas, although some of the bankfull discharge values fell within the range of error associated with the regression line. The difference between the bankfull discharge at each site and the 1.5-year discharge estimated from the WKCF regional regression for the same drainage areas was substantial: all 11 sites for which the values were compared had a Q_{bkf} less than or equal to 15% of the estimated regional $Q_{1.5}$; (Figure 4.5). Therefore, the regional 1.5-year discharge for WKCF streams, and estimates derived from the regression equations for Q_{bkf} would likely be more accurate than those derived from assumed flood frequencies. Furthermore, given the substantial difference between the compared values, bankfull return periods for each of the study sites could be expected to be more frequent than the approximately 1.5-year average reported for streams in other regions of the US (Dunne and Leopold 1978; Leopold et al. 1964; McCandless and Everett 2002; Mulvihill et al. 2005; Rosgen 1996; Williams 1978).



Figure 4.4 Regression curves for $Q_{1.01}$ and $Q_{1.5}$ derived from WKCF gauges. Bankfull flows estimated for study sites are plotted without regression lines.



Figure 4.5 Bankfull discharge as a proportion of the flow associated with the 1.5-year return interval for the same site. (Drainage area increases from left to right.)

5. Application of Bankfull Regional Relations

Regional curves describe characteristics that can generally be expected for streams of a given drainage area within a physiographic region. These descriptions are useful in the evaluation of stream stability, which includes the assessment of channel siltation, degradation, and bank erosion—factors that have substantial effects on aquatic habitat and sediment loads. They may be particularly useful in assessing channels undergoing rapid change, when bankfull indicators may be unapparent or ambiguous. Furthermore, these regional relations can be used as a basis for some restoration design methods (Rosgen 1998).

The regional curves for the WKCF region were developed from sites with watersheds between 0.25 and 117 mi² (from 1.56 to 117 mi² for silt-bed streams and from 0.25 to 2.88 mi² for gravel-bed streams) where the channel was stable relative to other streams of the region and had unambiguous bankfull indicators. The relationship between bankfull parameters and drainage area in the WKCF region is described by curves that explain between 54% and 92% of the variation within the datasets for bankfull area, width, depth, and discharge. Standard errors are less than 36%. Bankfull discharge was found to be no more than 15% of the estimated regional (WKCF) 1.5-year discharge and is thus likely associated with a smaller return interval. Therefore, use of the 1.5-year event for bankfull flow would represent a gross overestimate in all reaches examined in this project; estimates based on morphological features and/or the regional curves would be more accurate.

The curves developed in this project will be most applicable to streams having characteristics consistent with those criteria used to select the assessment reaches:

- Physiographic region. These curves apply to those streams with significant portions of their watersheds within the WKCF region.
- Land use. Streams in watersheds that are less than 10% urbanized are represented. Drainage projects and channelization have occurred in all of the watersheds in the region, and mining has occurred in several of the watersheds; therefore, the curves represent the effects of typical land use and sediment loads from channel alteration and mining.
- Flow regulation. Streams that are not subject to flow regulation are represented.
- Drainage area. The curves apply only to silt-bed streams draining between 1.56 and 117 mi² or gravel-bed streams draining between 0.25 and 2.88 mi².
- Sediment size. Silt- and gravel-bed streams are represented. Few sand-bed streams were located within the region, and none were found to meet the morphologic criteria for selection.
- Slope. The curves apply only to streams with slopes of up to 2%.

Streams affected by downstream confluences of large streams or locally high or large-caliber sediment supplies are not represented in the dataset used to develop these curves. Therefore, bankfull characteristics of channels formed under these conditions may be substantially different.

Because regional curves provide regional estimates of bankfull parameters that broadly describe stream conditions, they do not predict channel parameters for specific conditions that would form channels at specific sites. The cross-sectional dimensions of a channel are the product of many complex geomorphic processes, including the transport of sediment and channel evolution after repeated disturbance. A combination of geologic factors, the sequence and magnitude of land-use activities, and the sequence of channelization of the stream networks all have significant effects on sediment loads and channel evolution. Local watershed and channel conditions may cause channel bankfull flows and bankfull dimensions to differ significantly from those estimated from the equations produced by this project. Therefore, these equations should not be the only data used to evaluate or estimate bankfull characteristics in the assessment or design of WKCF channels. Rather, they should only be used in conjunction with field-based geomorphic assessment of the stream and its watershed. The results of field examination of bankfull conditions on the stream of interest should be compared to the WKCF regional curves. Channel dimensions that are more than one standard deviation greater or less than those dimensions estimated from the curves should be examined carefully to determine the cause of the variation. Likewise, designs that call for channel dimensions outside that range should provide sufficient data to justify the deviation from the curves.

Highly altered watershed conditions and direct manipulation of streams have changed watershed hydrology, sediment regimes, channel gradients, and base levels; ongoing maintenance continues to affect channel response and evolution. These altered reaches, from which the WKCF regional curves were developed, represent the geometry of evolving contemporary channels; if the channels were to completely recover from disturbance, their floodplains, planform patterns, and profiles would change, and their channel cross section characteristics would likely differ from those described by these regional curves. Therefore, if a restoration project intends to create bankfull characteristics similar to those that could be expected in a completely recovered channel, the design may require smaller dimensions than those that would be estimated from these curves.

References

- Amos. 1970. Blackford Quadrangle (O14).
- <u>ftp://ftp.kymartian.ky.gov/gqmaps1z/o14_gq.sid</u>, accessed Jun2009.
- Brush, LM. 1961. Drainage basins, channels and flow characteristics of selected streams in central Pennsylvania. USGS Professional Paper 282F:145–181.
- Burroughs, WG. 1924. The geography of the Western Kentucky Coal Field. Kentucky Geological Survey, Frankfort, KY, 205 pp.
- Dunne, T and LB Leopold. 1978. Water in environmental planning. WW Freeman and Co., New York, NY.
- Emmett, WW. 2004. A historical perspective on regional channel geometry curves. Stream Notes, Stream Systems Technology Center, USDA, Forest Service, Rocky Mountain Research Station, Fort Collins, CO, 2 pp. Jan.
- Franklin. 1969. Nebo Quadrangle (O16). Kentucky Environmental and Public Protection Cabinet, Office of Information, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/gqmaps1z/o16_gq.sid</u>, accessed Jun2009.
- Hansen and Smith. 1978. Livermore Quadrangle (O20). Kentucky Environmental and Public Protection Cabinet, Office of Information, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/</u> <u>gqmaps1z/o20 gq.sid</u>, accessed Jun2009.
- Harman, WA, GD Jennings, PM Patterson, DR Clinton, LO Slate, AJ Jessup, JR Everhart, and RE Smith. 1999. Bankfull hydraulic geometry relationships for North Carolina streams. AWRA Wildland Hydrology Symposium Proceedings, DS Olsen and JP Potyondy (eds.). AWRA Summer Symposium, Bozeman, MT.
- Harrelson, CC, CL Rawlins, and JP Potyondy. 1994. Stream channel reference sites: an illustrated guide to field technique. General Technical Report RM-245, US Department of Agriculture, Forest Service, Rocky Mountain Forest and Range Experiment Station, Fort Collins, CO, 61 pp.
- Johnson and Smith. 1972. Utica Quadrangle (N21). Kentucky Environmental and Public Protection Cabinet, Office of Information, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/</u> <u>gqmaps1z/n21 gq.sid</u>, accessed Jun2009.
- Johnson. 1973. Henderson Quadrangle (L17). Kentucky Environmental and Public Protection Cabinet, Office of Information, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/</u> <u>gqmaps1z/l17_gq.sid</u>, accessed Jun2009.

- Johnson. 1973. Wilson Quadrangle (L16). Kentucky Environmental and Public Protection Cabinet, Office of Information, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/gqmaps1z/l16_gq.sid</u>, accessed Jun2009.
- KDOW (Kentucky Division of Water). 2008. 2008 303(d) list of waters for Kentucky. Environmental and Public Protection Cabinet, Kentucky Division of Water, Frankfort, KY. Available at <u>http://www.water.ky.gov/sw/tmdl/303d.htm</u>, accessed Jun2009.
- KGS (Kentucky Geological Survey). 1980. Physiographic diagram of Kentucky. University of Kentucky, Lexington, KY. Scale not specified.
- KGS (Kentucky Geological Survey). 2002. Physiographic regions of Kentucky. Available at <u>http://kgsweb.uky.edu/download/state/regions.zip</u>, accessed 24Jun2005.
- Kilpatrick, FA and HH Barnes. 1964. Channel geometry in Piedmont streams as related to frequency of floods. USGS Professional Paper 422-E:10.
- Knighton, AD. 1987. River channel adjustment: the downstream dimension. In: River Channels: Environment and Process, KS Richards (ed.). Blackwell, Oxford, England, pp. 95-128.
- KYDGI (Kentucky Division of Geographic Information). 2005a. County Boundary Polygons of Kentucky. Kentucky Natural Resources Cabinet, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/county/County_Poly_1</u> z.ZIP, accessed Jun2009.
- KYDGI (Kentucky Division of Geographic Information). 2005b. National Hydrography Dataset (NHD) 24k streams of Kentucky (basin). Available at <u>ftp://ftp.kymartian.ky.gov/NHD/nhdorder/green6.e00</u> and <u>ftp://ftp.kymartian.ky.gov/ NHD/nhd-order/lowerohio6.e00</u>, accessed Jun2009.
- KYDGI (Kentucky Division of Geographic Information). 2006. Lower Ohio peak flows and Green River peak flows. Available at <u>ftp://ftp.kymartian.ky.gov/usgs/looh6pf.zip</u> and <u>ftp://ftp.kymartian.ky.gov/usgs/greenpf6.zip</u>, accessed Jun2009.
- Leopold, LB and T Maddock. 1953. The hydraulic geometry of stream channels and some physiographic implications. USGS Professional Paper 252, 57 pp.
- Leopold, LB, MG Wolman, and JP Miller. 1964. Fluvial processes in geomorphology. WH Freeman and Company, San Francisco, CA.
- McCandless, TL and RA Everett. 2002. Maryland stream survey: bankfull discharge and channel

characteristics of streams in the Piedmont hydrologic region. CBFO-S02-01, US Fish & Wildlife Service, Chesapeake Bay Field Office, Annapolis, MD, 40 pp.

McCuen, RH. 1998. Hydrologic analysis and design. Prentice Hall, Upper Saddle River, NJ.

McFarlan, AC. 1943. Geology of Kentucky. University of Kentcky/Waverly Press, Baltimore, MD, 531 pp.

McGrain, P. 1983. The Geologic Story of Kentucky. Special Publication 8, Series XI, Kentucky Geological Survey, Lexington, KY, 74 pp.

Mulvihill, CI, AG Ernst, and BP Baldigo. 2005. Regionalized equations for bankfull discharge and channel characteristics of streams in New York State: Hydrologic Region 6 in the southern tier of New York. USGS Scientific Investigations Report 2005-5100, 14 pp.

Noger, MC. 2002. Simplified geology of Kentucky, 1:500,000. Kentucky Geological Survey. Available at <u>http://kgsweb.uky.edu/download/state/</u> <u>kygeol.ZIP</u>, accessed Jun2005.

NRCS (National Resources Conservation Service). 2007. Regional hydraulic geometry curves. Available at <u>http://wmc.ar.nrcs.usda.gov/</u> <u>technical/HHSWR/Geomorphic/index.html</u>, accessed Mar 2007.

Palmer. 1968. Nortonville Quadrangle (Q18). Kentucky Environmental and Public Protection Cabinet, Office of Information, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/</u> <u>gqmaps1z/q18_gq.sid</u>, accessed Jun2009.

Palmer. 1972. Central City Quadrangle (P21). Kentucky Environmental and Public Protection Cabinet, Office of Information, Frankfort, KY. Available at <u>ftp://ftp.kymartian.ky.gov/gqmaps1z/</u> <u>p21_gq.sid</u>, accessed Jun2009.

Parola, AC, K Skinner, AL Wood-Curini, WS Vesely, C Hansen, and MS Jones. 2005. Bankfull characteristics of select streams in the Four Rivers and Upper Cumberland river basin management units. Project Final Report for Kentucky Division of Water NPS 99-12, University of Louisville Stream Institute, Louisville, KY, 30 pp.

Parola, AC, WS Vesely, WL Wood-Curini, DJ Hagerty, MN French, DK Thaemert, and MS Jones.
2005. Geomorphic characteristics of streams in the Mississippi Embayment physiographic region of Kentucky. Project Final Report for Kentucky Division of Water NPS 99-30, University of Louisville Stream Institute, Louisville, KY, 49 pp. Rice, CL. 2001. Pennsylvanian system. *In* The geology of Kentucky—a text to accompany the geologic map of Kentucky, RC McDowell (ed.) USGS Professional Paper 1151-H. Available at .<u>http://pubs.usgs.gov/pp/p1151h/penn.html</u>, accessed Mar2009.

Rosgen, DL. 1996. Applied river morphology. Wildland Hydrology, Fort Collins, CO.

Rosgen, DL. 1998. The reference reach—a blueprint for natural channel design. In: Proceedings of the 1998 ASCE Wetlands and River Restoration Conference, Denver, CO. March.

Smith, DP and L Turrini-Smith. 1999. Western Tennessee fluvial geomorphic regional curves. Submitted to the US EPA Region IV as part of requirements for the grant associated with agreement #CP984142-97-4.

Strahler, AN. 1957. Quantitative analysis of watershed geomorphology. American Geophysical Union Transactions 38:913-920.

Tasker, GD. 1978. Relation between standard errors in log units and standard errors in percent. WRD Bulletin, Jan-Mar – Apr-June:86–87.

USEPA (US Environmental Protection Agency). 1999. Protocol for developing sediment TMDLs. EPA 841-B-99-004. Office of Water (4503F), Washington, DC, 132 pp. October.

USEPA and USGS. 2005. Strahler stream order. National Hydrography Dataset Plus (NHDPlus). Available at <u>ftp://ftp.horizon-systems.com/</u> <u>NHDPlusExtensions/SOSC/NHDPlus05V01_01</u> <u>sosc.zip</u>, accessed Oct2008.

USIACWD (US Interagency Advisory Committee on Water Data). 1982. Guidelines for determining flood flow frequency. Bulletin 17-B of the Hydrology Subcommittee, USGS Office of Water Data Coordination, Reston, VA, 183 pp.

Williams, GP. 1978. Bank-full discharge of rivers. Water Resources Research 14(6):1141–1153.

Wolman, MG. 1955. The natural channel of Brandywine Creek, Pennsylvania, US. USGS Professional Paper 271:87–109.

Wolman, MG and LB Leopold. 1957. River flood plains: some observations on their formation. United States Geological Professional Paper 282B:87-110.

Wolman, MG and JP Miller. 1960. Magnitude and frequency of forces in geomorphic processes. Journal of Geology 68:54–74.

Appendices

Financial and Administrative Closeout

А

PROJECT OUTPUTS

Mi	lestones	Expected Begin Date	Expected End Date	Actual Begin Date	Actual End Date
1.	Submit all draft materials to EPPC for review and approval.	Duration			
2.	Submit advanced written notice of all workshops, demonstrations, and/or field days to EPPC.	Duration			
3.	QA/QC submission and approval.	Jul 04	Jul 04	Dec 02	Jan 05
4.	Develop gauging station list.	Nov 04	Jun 05	Oct 05	Dec 05
5.	Develop reference reach list.	Dec 04	Jun 07	Oct 05	Mar 08
6.	Analyze reference site frequency.	Dec 04	Feb 07	Oct 05	Mar 08
7.	Collect reference site data.	Jun 05	Dec 07	Mar 07	Apr 09
8.	Analyze data from reference reaches.	Nov 05	Jan 08	Mar 07	Apr 09
9.	Develop regional curves.	Aug 06	Sept 08	Jul 07	Apr 09
10.	Upon request, submit annual report and/or participate in the EPPC-sponsored biennial NPS Conference.	Duration			
11.	Submit three copies of the final report and submit three copies of all products of this project.	Sept 08	Mar 09		Jun 09

Non-Federal **Budget Categories** Section 319(h) Match Total \$ Personnel \$ 73,723 \$ 30,380 104,103 Supplies 2,000 0 2,000 Equipment 2,000 12,000 14,000 Travel 7,025 7,025 0 Contractual 0 0 0 Operating 22,128 33,180 55,308 Other 6,200 6,200 0 Total \$ 113,076 \$ 75,560 \$ 188,636

BUDGET SUMMARY

Budget Categories	Sect	ion 310(h)	Noi	1-Federal Match	Total	Fv	Final nonditures	T	nenant
Budget Categories	Sect	ion 319(n)	-	viaun	10141	ĽA	penuntui es	U	nspent
Personnel	\$	79,573	\$	29,240	\$ 108,813	\$	105,985	\$	2,828
Supplies		2,000		0	2,000		4,165		(2165)
Equipment		860		13,140	14,000		11,279		2,721
Travel		2,705		0	2,705		1,796		909
Contractual		0		0	0		0		0
Operating		21,738		33,180	54,918		53,768		1,150
Other		6,200		0	6,200		6,415		(215)
Total	\$	113,076	\$	75,560	\$ 188,636	\$	183,408	\$	5,228

KDOW APPROVED REVISED BUDGET

The University of Louisville Research Foundation was reimbursed \$109,935.02. A total of \$3,140.98 federal funds remain unspent. These excess funds result primarily from final personnel and supplies costs being lower than expected.

EQUIPMENT SUMMARY

Type of Equipment	Estimated Cost	Actual Cost	Balance	
Lease of four-wheel-drive field research vehicle	\$ 6,000.00	\$ 4,941.18	\$ 1,058.82	Rates per common and contracted UofL vendors
Maintenance of field research vehicle	1,000.00	66.03	933.97	Below estimate
Maintenance of robotic total station	1,000.00		1,000.00	Not needed
Laptop computer and accessories	6,000.00	6,271.51	(271.51)	Professional tablet PC
Total	\$ 14,000.00	\$ 11,278.72	\$ 2,721.28	

None of the equipment purchased has a current fair market value exceeding \$5,000.

SPECIAL GRANT CONDITIONS

A Quality Assurance Project Plan (see Appendix B) was submitted in December 2002.

Quality Assurance Project Plan **B**

Prepared by: Arthur C. Parola, Jr., Ph.D. Department of Civil and Environmental Engineering University of Louisville

April 2001

PROJECT AND QA MANAGER: Arthur C. Parola, Jr., Ph.D. Professor of Civil and Environmental Engineering University of Louisville Louisville, KY 40292 <u>a.c.parola@louisville.edu</u> (Phone) 502-852-4599 (Fax) 502-852-8851

Group A: Project Management Elements

A3 Distribution List

Mrs. Margi Jones Riparian Management/Restoration Advisor Kentucky Division of Water 14 Reilly Rd. Frankfort, KY 40601 502/564-3410 502/564-0111 (Fax)

A4 Project /Task Organization

The stream geomorphic data collected in this project will be used by individuals assessing stream stability in the specific region characterized by the stream geomorphic data. The data will also be used by designers of stream restorations to determine the likely range of stream characteristics of the proposed stream restoration.

The project QA manager will be Dr. Parola. A research associate with training in applied geomorphology will collect all field data with the assistance of graduate students and professional staff. Dr. Parola, the project director, will maintain the official, approved QA Project Plan.

Figure B.1 illustrates the relationships and lines of communication between all project participants.



Figure B.1 Organizational Chart Showing Lines of Communication.

A5 Problem Definition and Background

Stream physical habitat, stream stability, bank erosion and total sediment loads are affected by the physical characteristics or stream channel networks of a watershed. Land-use practices in Kentucky, such as land development, livestock grazing, land clearing, channel relocation and modifications for flood protection, roadway construction and mining tend to increase stream peak flow rates, disturb riparian vegetation and alter stream channel characteristics. The response of many streams to disturbance can be excessive production of sediments through channel incision followed by severe bank erosion. In many cases in Kentucky, channels have incised into bedrock and have continued to widen through bank erosion for decades after disturbances have occurred. The direct disturbances to streams and the associated indirect erosion that continues for long periods can severely degrade stream habitat upstream, downstream and at the disturbed section of the stream. Physical alterations of stream channels are a significant source of stream habitat degradation and a major source of non-point source pollution in Kentucky's watersheds.

Methods of stream restoration and bioengineering techniques have been developed to improve stream habitat, reduce bank erosion and reduce sediment loads through physical alteration of disturbed stream channels. Determination of the necessity for stream restoration and restoration design requires that stream physical characteristics be compared to data from stable reference reaches in the same climactic and geophysical regions. Regional data on geomorphologic characteristics of streams are an important source for information necessary for stream evaluation and restoration design. At present, stream restoration in Kentucky is being conducted without the benefit of regional information on geomorphic parameters, although data collection in one River Basin Management Unit (RBMU) has been completed and the data has been analyzed. Data of bankfull characteristics for streams in the Tennessee/Cumberland/Mississippi RBMU were collected in 2000 (FFY 1999). The project will extend the collection, analysis and development of regional geomorphic stream characteristic curves to physiographic regions in which data have not been collected.

A6 Project Task Description

Collect Available Site Information

Information on all stream gauge stations in the selected physiographic region will be obtained to determine the possible locations of assessment sites. This information will be used to develop a list of potential sites for data collection.

Select Assessment Reaches

Each of the stream gauge stations selected as possible data collection sites will be visited. A preliminary classification of the stream type (Rosgen 1996) will be made. Sites that have ambiguous bankfull characteristics or other characteristics that make them unsuitable for use as assessment sites will be eliminated. The remaining stream gauge station sites will be considered for further analysis.

Experience in collection of data in the Tennessee/Mississippi/Cumberland RBMU has shown that channel conditions at stream gauge stations are typically significantly different than those upstream or downstream of stream gauge stations where channels have developed bankfull indicators. Debris collection, channel maintenance and bridge construction at or near stream gauge stations obscure or prevent the formation of a stable channel configuration; consequently, the information from the stream-gauge stations may not be valuable at most or all stream gauges. If similar problems are found at stream gauge stations in this project area, assessment reaches away from stream gauge stations will be selected. Extensive reconnaissance will be required to locate stable channels without stream gauge stations.

Complete Analysis of Assessment Site Station Frequency

Analysis of peak flow frequency will be conducted to develop a flood frequency curve from which the 1-to-2-year flow event and water surface elevation can be determined.

Collect Data of Assessment Reaches for Geomorphic Parameters

The site data required to characterize the bankfull conditions will be collected at each assessment site. The information will be collected to describe channel geomorphic characteristics over riffle sections as described in Rosgen (1996) for characterization of assessment reaches. Data sheets and photographs of each site will be developed to be useful to others conducting restoration work in the area of the assessment sites. The information will be stored and made available as part of a spreadsheet database. The Nonpoint Source Section of the Kentucky Division of Water will receive an electronic copy of the data. The information will be stored in a format that will be transferable to the Kentucky Division of Water for conducting watershed evaluation, restoration, or TMDL projects.

Where assessment reach information must be collected at sites without stream gauge station information, cross section data will be collected over riffle zones to the extent necessary to develop and run models of flow (using HEC-RAS) to approximate a range of channel formative discharges.

Analyze Bankfull Characteristics of Assessment Reaches

The geomorphic data collected from each assessment site will be analyzed. A definitive stream classification based on the site measurements will be made. Bankfull flow rates will be extracted or modeled from bankfull elevation measurements and compared with frequency analysis information. Roughness coefficients will be computed from the cross section, bankfull field information, and stream gauged flow rates where stream gauge information is available. Average velocities, depths, boundary stresses and stream power of bankfull flow will be computed.

Develop Regional Data and Curves of Bankfull Geomorphologic Characteristics

Data representing bankfull geomorphologic characteristics and variability will be developed for each physiographic region. Data for each assessment reach will be displayed according to stream type. The information will be presented in a clear and simple format.

Submit Annual, Final and Closeout Reports

An annual report will be submitted. The University of Louisville Research Foundation will request current final project and closeout report guidelines from the Kentucky Division of Water no less than six months prior to the project end date. The final project report will present the data and analysis of each assessment site in a clear and standard format. The data from each of the assessment sites will be stored in a database. The Nonpoint Source Section of the Kentucky Division of Water will receive an electronic copy of the database. The report will also present and describe the regional data that represent bankfull geomorphic characteristics. A closeout report will be prepared and submitted as required by the US Environmental Protection Agency.

A7 Quality Objectives and Criteria

The objective of this project is to develop reliable regional bankfull characteristics of stream channels in a specific physiographic region of Kentucky. Two basic groups of data will be collected: sediment samples and stream geometric characteristics. In addition, hydrologic data from USGS gauging stations will be used to associate flows with specific recurrence intervals.

Surveying techniques that provide accuracy of about 1 cm in all directions will be used with the total station equipment that will be employed for stream geometric data collection. Also standard sieve analysis procedures employed by the geomechanics laboratory using standard ASTM techniques for fine and coarse aggregates will provide data for sediment size gradation to high precision. Large variations in geometric characteristics (typically on the order of 0.3 m) are associated with the subjective selection of bankfull elevations based on field indicators; therefore all bankfull indicators will be measured and flow levels associated with each indicator will be reported. These indicators include tops of coarse bar deposits, tops of fine bar deposits, low vegetation lines, tops of banks and floodplain elevations.

Sediment sampling in coarse bed channels is limited by the ability to only sample a very small portion of the streambed. Four techniques may be used to assess sediment in gravel and cobble bed streams:

- 1. Pebble counts on each riffle studied
- 2. Riffle subsurface bulk samples
- 3. Bar bulk samples
- 4. 30 largest particles on the bar

Amounts of gravel required to characterize the active streambed will be determined according to Bunte and Abt (2001), Rosgen (1996) and Kappesser (2002).

To ensure consistency in the selection of sampling locations for bankfull indicators, for collection of geometric stream characteristics, and for sampling of bar materials, the QA manager will conduct on-site quality checks.

The USGS maintains well-established quality control procedures for the gauge data flows. The quality of each measurement is recorded. Only good or excellent measures of flow will be used in the assessment of bankfull flow conditions.

A8 Special Training and Certification

The QA manager and project team have acedemic as well as professional training in applied morphology and the techniques necessary to collect and analyze the required geomorphic data. This training includes extensive academic and professional training in surveying, sediment sampling, hydraulic and hydrologic modeling, and geomorphic assessment.

A9 Documents and Records

The QA manager will be responsible for ensuring that the data collection team and all others on the distribution list have the most current QA project plan through email distribution. However, we do not anticipate significant changes to the QAPP. The data report will include the items described in Table B.1.

A final report that documents the procedures used to collect the data, difficulties in the data collection process and factors that influenced data quality will be produced. The final report will include the raw field data in a database. The final report will also include the analysis and products derived from the analysis such as regional curves of stream geomorphic characteristics and any other relations derived from these analyses. The database will be submitted to the Kentucky Division of Water project manager.

Group B: Data Generation and Acquisition Elements

B1 Sampling Process Design (Experimental Design)

Regional curves have been developed using USGS stream gauge station information. Stream geometric properties near the gauge stations have been used to determine bankfull characteristics, flow rates and flow frequency. Ideally, bankfull curves can be developed solely from information at USGS gauge stations within a physiographic region of Kentucky.

Previous work at gauge stations in the Cumberland and Tennessee RBMU demonstrated that most sites lacked consistent and reliable bankfull indicators because of the frequent disturbance near the gauge station. Gauge stations were typically located on or near bridges that frequently accumulated debris and severe bank erosion occurred around the bridge. Alternatively, channels were heavily straightened and lacked benches or other well-defined bankfull features. In these watersheds and on streams with slopes greater than 0.5%, reference reaches were found with reliable bankfull indicators and sufficient geometric data were obtained to model flows through relatively straight riffle sections. Modeling was considered unreliable on streams with smaller slopes and only cross section geometric data were obtained.

Type of Data	Source Data	Analysis
Site location, geology and topographic data	USGS topo maps, KY geologic maps, state GIS database	None
Bankfull stream characteristics	Geometric data collected by field team	Use of HEC-RAS flow model and AutoCad
Sediment gradation characteristics	Sediment data collected by field team	Grain size analysis
Streamflow and frequency distribution	USGS streamflow gauge data	Peak flow frequency modeling techniques

Table B.1Final Report Data

In consideration of the above information, one or a combination of the following methods will be used:

- 1. Obtain stable and reliable bankfull indicators at gauge station;
- 2. Model flows on streams with sufficient slope (greater than 0.5%); and
- 3. Obtain only geomotric characteristics of channel on low-gradient streams.

The procedures for these methods are outlined in Rosgen (1996); the flow modeling is included in Brunner (2001).

B2 Sampling Methods

Sampling for this project can be grouped into two categories: (1) surveying for channel geometric characteristics and (2) sediment sampling. Table B.2 describes the types of data to be sampled and the method used to sample.

Survey data will be checked during the surveying process by intermittently checking elevations at monumented locations. Any error in survey information will be apparent by following standard professional surveying procedures. A resurvey will be initiated when errors occur.

Total sediment weight before and after sieve analysis will be used to determine the error in sieve analysis procedures. Samples with an error greater than 8% will not be used, and the reasons for the errors will be determined and corrective action will be taken. The QA manager will be responsible for reviewing the sediment grain size distribution error analysis to determine the need to repeat the analysis.

Survey errors are most often apparent in the field when control points are recorded. Maximum errors at control points will be recorded. Surveys will be repeated where the errors at monuments are greater than 2 cm. The QA manager will review survey error measures at each site to ensure that inaccurate surveys are repeated.

B3 Sample Handeling and Custody

Total station survey data will be collected in electronic format on data loggers and downloaded each day to a laptop computer. Pebble count and other sediment data will be recorded on data forms and typed into a database.

Sediment samples will be labeled in the field and transported directly to the geomechanics laboratory. Grain size analysis will be conducted in the laboratory within one month of sample collection. Grain size analysis will be completed and data will be directly entered into a computer database. The data will be archived by the project QA manager.

Type of Data	Method	Reference
Channel cross section	Total station survey	Rosgen 1996
Channel profile	Total station survey	Rosgen 1996
Channel planform	Total station survey	Rosgen 1996
Riffle surface sediment grain size distribution	Wolman pebble counting	Bunte and Abt 2001
Subsurface sediment grain size distribution	Fine and coarse sieve analysis	Bunte and Abt 2001
Bar sediment grain size distribution	Fine and coarse sieve analysis	Rosgen 1996; Bunte and Abt 2001
Largest particles on bar size distribution	Field measured using ruler	Kappesser 2002

Table B.2 Sampling Methods

B4 Analytical Methods

Survey data will be analyzed and reduced using AutoCad. Cross section and stream profile characteristics will be extracted from the AutoCad data for further analysis using Microsoft Excel. The data will be entered into a Microsoft Access database following quality control checks during data processing and confirmation of satisfactory quality through spreadsheet analysis.

Gauge data frequency analysis will be conducted using several hydrologic modeling techniques. Peak flow estimates of flow frequency are unreliable at the level of channel formative and bankfull flow conditions. Alternative methods for quantifying flow frequency are being investigated.

The open channel flow model HEC-RAS will be used to obtained bankfull flow rates.

Linear regression techniques will be used to obtain regional relations for bankfull geomorphic parameters.

B5 Quality Control

Bulk sediment sample weights will be compared before and after sieve analysis to determine the percentage lost in the sieving processes. A loss of less that 8 % will be considered adequate for the sampling required for characterizing the bed sediments.

Standard surveying practices will be employed to ensure that survey location error is less than 1 cm.

B6 Instrument and Equipment Testing, Inspection, and Maintenance

Survey equipment and scales will be maintained to ensure proper function. This equipment will be tested against standards before and after field reconnaissance. The equipment will be sent to a local survey company for recalibration if found to be inaccurate or out of calibration

Sieves are cleaned after each use. Damaged sieves will be replaced.

B7 Instrument and Equipment Calibration and Frequency

Survey equipment is calibrated every six months, although it may be calibrated more frequently if found to be out of calibration during testing.

B8 Inspection and Acceptance of Supplies and Consumables

This does not apply to geomorphic data collection.

B9 Non-direct Measurements

Annual peak flows and gauge station rating curves will be obtained from the USGS. Strict and rigouous QA and QC has been established by the USGS to ensure the quality of these data. Ratings are given to flow data such that measurements of rating less than good will not be accepted for use in this project.

B10 Data Management

The data will be archived in paper format and entered into an Excel spreadsheet and archived.

Group C: Assessment and Oversight Elements

C1 Assessment and Response Actions

Assessment of data quality will be conducted at several levels. Survey equipment will be examined to determine its accuracy by laying out a known measurement distance and through repeat measurements each time the equipment is taken into the field.

The QA data manager will make visits to field sites during part of each field reconaissance to ensure that procedures described here are being followed. The project team will discuss procedures and assess errors in measurements at least biannually. Data collection will be repreated if necessary.

Accuracy of the surveying equipment is imperative for high quality field measurements. At least one backup instrument will be made available to ensure that a calibrated instrument is used.

C2 Reports to Management

Verbal reports on the status of projects will be made weekly. Data collecton procedures will be discussed, problems will be addressed and any necessary corrective actions will be taken on a weekly basis. The QA manager and field data collection team will meet to discuss QA and QC issues before each intenseve field data collection period.

Group D: Data Validation and Usability Elements

D1 Data Review, Verification and Validation

Spot checks of data using a simple level line and tape will be made to ensure that survey data are within an acceptable range for characterizing geomorphic parameters. Most problems with data error will be addressed at the time of data collection.

D2 Verification and Validation Methods

The geomorphic data and regional relations for this project will be compared to those of other similar projects of regional geomorphic characteristics such as those by McCandless and Everett (2002). Data incorporated in the database will be reviewed and tested by the QA manager. Although large variationin geomorphic paramaters is anticipated; unusal deviation will be examined carefully to ensure that the they represents variation in geomorphic characteristics and not error of data collection and analysis procedures.

D3 Reconciliation and User Requirements

The antipated users of this data and the resulting regional relations developed from the data are individuals conducted stream assessments or designers of stream restoations. Large natural variations of sediment supply and stream geomorphic characteristics occur because of variation in current landues, landuse history, and direct modification of satream channels as well as variation in geology. Users of the information will be warned of that the data may be biased toward streams on which gauge stations have been installed (larger watersheds). In addition, users will be warned that local and basin conditions may cause substatial diffenences in stream characteristics. The database will provide information on the characteristics of streams and their watersheds, so that users have information available to make direct comparisons with specific site conditions.

References

Andrews, ED. 1980. Effective and bankfull discharges of streams in the Yampa River Basin, Colorado and Wyoming. Journal of Hydrology 46:311-330.

Andrews, ED. 1984. Bed material entrainment and hydraulic geometry of gravel bed rivers in Colorado. Bulletin of the Geological Society of America 95:371.

Andrews, ED. 1994. Marginal bedload transport in a gravel bed stream, Sagehen Creek, California. Water Resources Research 30(7):2241-2250.

Andrews, ED and DC Erman. 1986. Persistence in the size distribution of surficial bed material during an extreme snowmelt flood. Water Resources Research 22(2):191-197.

Brunner, G. 2001. HEC-RAS, river analysis system, hydraulic reference manual, Version 3.0," US Army Corps of Engineers Report No. CPD-69.

Bunte, K and SR Abt. 2001. Sampling surface and subsurface particle-size distributions in wadable gravel-and cobble-bed streams for analyses in sediment transport, hydraulics, and streambed monitoring. Gen. Tech. Rep. RMRS-GTR-74, USDA Forest Service, Rocky Mountain Research Station, Fort Collins, CO, 428 p.

FISRWG (Federal Interagency Stream Restoration Working Group). 1998. Stream corridor restoration: principles, processes, and practices. GPO Item No. 0120-A; SuDocs No. A 57.6/2:EN 3/PT.653, FISRWG (15 Federal agencies of the US Govt.). CD-ROM.

Kappesser, GB. 2002. A riffle stability index to evaluate sediment loading to streams. Journal of the American Water Resources Association (38)4:1069-1081.

Knighton, D. 1998. Fluvial forms and processes, a new perspective. Edward Arnold Publishers, London, UK. Leopold, LB and T Maddock, Jr. 1953. The hydraulic geometry of stream channels and some physiographic implications. USGS Professional Paper 252, 57 pp.

Limerinos, JT. 1970. Determination of the Manning roughness coefficient for measured bed roughness in natural channels. USGS Water Supply Paper 1898B, 47 pp.

McCandless, TL and RA Everett. 2002. Maryland stream survey: bankfull discharge and channel characteristics of streams in the Piedmont hydrologic region. CBFO-S02-0, US Fish & Wildlife Service, Chesapeake Bay Field Office.

Montgomery, DR and JM Buffington. 1993. Channel classification, prediction of channel response and assessment of channel condition. Report TFW-SHIO-93-002, Department of Geological Sciences and Quaternary Research Center, University of Washington, Seattle, WA.

Rosgen, DL. 1996. Applied river morphology, 2e. Wildland Hydrology, Pagosa Springs, CO.

- Schumm, SA. 1960. The shape of alluvial channels in relation to sediment type. USGS Professional Paper 352B:17–30.
- Schumm, SA, MD Harvey, and CC Watson. 1984. Incised channels: morphology, dynamics and control. Water Resources Publications, Littleton, CO.

Simon, A. 1989. A model of channel response in disturbed alluvial channels. Earth Surface Processes and Landforms 14:11-26.

Strahler, AN. 1957. Quantitative analysis of watershed geomorphology. American Geophysical Union Transactions 38:913-920.

Williams, GP. 1978. Bank-full discharge of rivers. Water Resources Research 14(6):1141-1154.

Western Kentucky Coal Field Data

The data reported for each assessment site are provided on the enclosed CD-ROM.

	USGS Gage			
Stream Name	Number	State	County	Latitude
Caney Creek*		KY	Ohio	N37.460317
Drakes Creek		KY	Hopkins	N37.174000
Drakes Creek at Old Nortonville-Whiteplains Road		KY	Hopkins	N37.180783
Eagle Creek Tributary		KY	Union	N37.637350
East Fork of Flynn Fork		KY	Caldwell	N37.140817
Hazel Creek*		KY	Muhlenberg	N37.147150
Lewis Creek		KY	Ohio	N37.350500
Lick Creek		KY	Hopkins	N37.268317
Lick Creek at Paul Peyton Rd		KY	Hopkins	N37.284317
Muddy Creek*		KY	Butler	N37.169600
No Creek*		KY	Rough	N37.487933
Otter Creek		KY	Hopkins	N37.489533
Pup Creek*		KY	Daviess	N37.842100
Slover Ditch Tributary*		KY	Webster	N37.447383
Welch Creek*		KY	Butler	N37.261217
West Fork Adams		KY	Hancock	N37.688833
West Fork Adams near Newton Springs Church		KY	Hancock	N37.682867
West Fork Adams Tributary		KY	Hancock	N37.687983
West Fork Pond River*		KY	Christian	N37.108244
Whitelick Creek***		KY	Henderson	N37.639967

Notes:

* - Remnant channel

** - Probable stream type before abandonment or incision

*** - Recently incised; only bankfull width measured

- + Interpreted from USGS topographic maps
- + Unable to determine

⁺ + Slope of sinuous valley incised within floodplain of larger stream

	Drainage	Date of	Bankfull Discharge	Bankfull Velocity	Bankfull
Longitude	Area (mi ²) Physiographic Region	Study	(CFS)	(ft/s)	Area (ft ²)
W86.653767	117.00 Western KY Coal Field	Mar-07	356	1.5	237.3
W87.444083	29.10 Western KY Coal Field	Mar-07	325	1.5	216.4
W87.435267	31.79 Western KY Coal Field	Mar-07			
W87.903333	0.77 Western KY Coal Field	Apr-07	25	3.0	8.3
W87.758800	2.88 Western KY Coal Field	Mar-07	43	3.0	14.2
W86.978667	10.97 Western KY Coal Field	Mar-07			
W86.912117	8.30 Western KY Coal Field	Mar-07	86	1.5	57.5
W87.714917	19.66 Western KY Coal Field	Mar-07	107	1.5	71.0
W87.723033	23.00 Western KY Coal Field	Mar-07	190	1.5	126.4
W86.772783	83.15 Western KY Coal Field	Mar-07			
W86.987983	9.40 Western KY Coal Field	Mar-07			
W87.384417	40.36 Western KY Coal Field	Mar-07	221	1.5	147.4
W86.968533	26.13 Western KY Coal Field	Apr-07			
W87.729333	1.56 Western KY Coal Field	Mar-07			
W86.595100	45.47 Western KY Coal Field	Mar-07			
W86.690250	0.87 Western KY Coal Field	Mar-07	27	3.0	9.1
W86.694717	1.53 Western KY Coal Field	Mar-07	44	3.0	14.8
W86.692667	0.25 Western KY Coal Field	Mar-07	11	3.0	3.6
W87.363503	43.00 Western KY Coal Field	Mar-07			
W87.700800	3.82 Western KY Coal Field	Apr-07			

Bankfull Width (feet)	Bankfull Depth (feet)	Maximum Bankfull Depth (feet)	Bankfull Hydraulic Radius (feet)	Bankfull Gage Height (feet)	W/D Ratio	Floodprone Width (feet)	Entrenchment Ratio
46.5	. 5.1	6.6	4.7		9.1	>102	>2.2
39.0	5.6	8.9	4.9		7.0	>86	>2.2
31.8							
12.8	0.7	1.0	0.6		19.7	18	1.4
13.6	1.0	1.6	1.0		13.1	22	1.6
17.7							
23.4	2.5	4.5	2.18		9.6	>51	>2.2
32.4	2.2	3.4	2.12		14.8	>71	>2.2
36.1	3.5	5.0	3.29		10.3	>79	>2.2
35.1							
36.4							
31.5	4.7	7.8	4.01		6.7	287	9.1
39.3							
13.0							
32.0							
10.9	0.8	1.1	0.79		13.2	17	1.6
16.2	0.9	1.8	0.87		17.7	23	1.4
6.0	0.6	1.0	0.5		10.0	14	2.4
48.0							
23.3							

	Riffle	Riffle			Bulk Bar		
Valley Incision	d50	d84	Manning	Bulk Bar	dmax	Bankfull	Bankfull Shear
Ratio (ft/ft)	(mm)	(mm)	n-value	d50 (mm)	(mm)	Slope (ft/ft)	Stress (lb/ft ²)
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
5.7	Gravel	Gravel					
2.8	Gravel	Gravel					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
1.0	Silt	Silt					
3.5	Gravel	Gravel					
3.2	Gravel	Gravel					
3.5	Gravel	Gravel					
1.0	Silt	Silt					
	Silt	Silt					

		Rosgen		1.5-year	Bedrock
Local	Local Valley	Stream	Bankfull Return	Discharge	Influence
Sinuosity	Slope (ft/ft)+	Туре	Period (years)	(CFS)	(yes/no)
2.3	0.0006	E6			Ν
1.6	0.0006	E6			Ν
1.6	0.0006	E6			Ν
1.1	0.0059	C4			Y
1.4	0.0016	B4c			Ν
1.6	0.0009	E6**			Ν
1.4	0.0010	E6			Ν
1.7	0.0011	E6			N
1.8	0.0010	E6			Ν
2.2	0.0003	E6**			Ν
2.1	†	E6**			Ν
1.8	0.0006^{\ddagger}	E6			Ν
1.6	0.0006	E6**			Ν
1.9	0.0008	E6**			Ν
2.0	0.0007	E6**			Ν
1.1	0.0115	B4c			Ν
1.1	0.0054	B4c			Ν
1.0	0.0175	C4			Y
2.3	0.0011	E6**			Ν
1.4	0.0018	E6**			Ν